Disproportionation of (Mg,Fe)SiO$_3$ perovskite in Earth's deep lower mantle
Li Zhang et al.
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while children had detectable antibodies to PfSEA-1. These data suggest that anti--PfSEA-1 responses could augment responses to other targets, such as PfEMP-1, that may protect against severe malaria (24). Under conditions of natural exposure, only 6% of 1.5- to 3.5-year-old children in our cohort had detectable antibodies to PfSEA-1. This low prevalence suggests that adventuated vaccination with PfSEA-1 could have a marked impact on reducing severe malaria in young children.

In a second longitudinal Kenyan cohort, anti--PfSEA-1 antibodies were associated with significant protection against parasitemia in adolescents and young adults. Individuals with detectable IgG anti--rPfSEA-1A antibodies had 50% lower parasite densities over 18 weeks of follow-up (as a candidate for pediatric falciparum malaria. – tentatively higher than in our birth cohort, which suggests that natural exposure boosts anti--PfSEA-1 antibody levels. Together, our data validate our field-to-lab-to-field-based strategy for the rational identification of vaccine candidates and support PfSEA-1 as a candidate for pediatric falci parum malaria. By blocking schizont egress, PfSEA-1 may synergize with other vaccines targeting hepatocyt e (25) and RBC invasion (26).

REFERENCES AND NOTES

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SUPPLEMENTARY MATERIALS
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30 October 2012; accepted 25 April 2014
10.1126/science.1254417

DEEP EARTH
Disproportionation of (Mg,Fe)SiO3 perovskite in Earth’s lower mantle
Li Zhang,1,2 Yue Meng,3 Wenge Yang,1,4 Lin Wang,1,4 Wendy L. Mao,5,6 Qiao-Shi Zeng,5 Jong Seok Jeong,7 Andrew J. Wagner,7 K. Andre Mkhoyan,7 Wenjun Liu,4 Ruqing Xu,4 Ho-kwang Mao1

The mineralogical constitution of the Earth’s mantle dictates the geophysical and geochemical properties of this region. Previous models of a perovskite-dominated lower mantle have been built on the assumption that the entire lower mantle down to the top of the D" layer contains ferromagnesiam silicate [(Mg,Fe)SiO3] with nominally 10 mol% Fe. On the basis of experiments in laser-heated diamond anvil cells, at pressures of 95 to 101 gigapascals and temperatures of 2200 to 2400 kelvin, we found that such perovskite is unstable: it loses its Fe and disproportionates to a nearly Fe-free MgSiO3 perovskite phase and an Fe-rich phase with a hexagonal structure. This observation has implications for enigmatic seismic features beyond ~2000 kilometers depth and suggests that the lower mantle may contain previously unidentified major phases.

The conventional view of materials in Earth’s lower mantle, which comprises the largest fraction of our planet (>55% by volume), has evolved greatly in the past 10 years. High-pressure-temperature (P-T) experiments of the past century suggested that the predominant phase was ferromagnesiam silicate [(Mg,Fe)SiO3] with the orthorhombic perovskite (pv) structure, unchanging over the enormous P-T range from the bottom of the transition zone at 670 km depth (24 GPa) to the core-mantle boundary (CMB) at 2900 km (326 GPa). However, laser-heated diamond anvil cell (DAC) technology coupled with x-ray spectroscopy (XRS) and x-ray diffraction (XRD) probes have led to two key discoveries: namely, the phase transition of (Mg,Fe)SiO3 pv to the post-perovskite (ppv) structure (1–3) near the D" layer of the lower mantle, and the pressure-induced spin-pairing of the 3d electrons in Fe(4–6)

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that dramatically alters the chemical and physical behavior of the Fe-bearing silicates and oxides (7). These discoveries have triggered a flurry of new theories for the enigmatic seismic and geodynamic features observed in the deep lower mantle, such as the geothermal and seismic signatures of the spin transition in Fe-bearing minerals, the correlations between the pv-ppv P-T boundary and the D″ discontinuity and velocity gradients, and possible mineralogical mechanism for the existence of large low-shear-velocity provinces (LLSVs) (1, 5, 8–10).

Recent sound-velocity measurements of silicate pv and (Mg,Fe)O ferropericlase (fp) suggested that the (Mg,Fe)SiO₃ pv–dominant petrological model (22) would be consistent with the density profile of the lower mantle (13). The MgSiO₃–FeSiO₃ pseudo-binary phase relations, however, have only been well investigated at the P-T conditions corresponding to the top half of the lower mantle (14, 15). Beyond 60 GPa, detailed studies of this phase diagram are still scarce and reconnaissance in nature (16–19).

**Disproportionation of (Mg,Fe)SiO₃**

To more directly interrogate these phase relations, we brought crystalline (Mg₀.₈₅Fe₀.₁₅)SiO₃ (Fs₁₅) orthopyroxene (opx) sample to the pressures (60 to 110 GPa) and temperatures (2000 to 2400 K) of the lower mantle using DAC technology (20). We monitored the growth of crystalline phases with in situ XRD under high P-T laser-heating conditions as well as after T-quench at high P.

Compressing Fs₁₅ to 95 GPa at 300 K, the opx powder XRD peaks broaden substantially, but the sample was not amorphized. Instead, the broad peaks resemble the major peaks of pv (fig. S4), indicating a nanocrystalline pv-like structure. The pattern sharpens with moderate heating at 1500 to 2000 K and forms normal crystalline pv, with unit cell parameters consistent with the 15% Fs composition (fig. S4). Increasing temperature to 2300 K above 95 GPa (fig. S4) or heating directly to this P-T condition (Fig. 1A), the pv–dominated powder XRD pattern exhibited subtle additional features. At 101 GPa, the pv unit cell of \( a = 4.363(3) \text{ Å} \), \( b = 4.611(4) \text{ Å} \), \( c = 6.349(3) \text{ Å} \), and molecular volume \( V = 31.89(3) \text{ Å}^3 \) (Fig. 1A), which is 7.8° in (B), belongs to a small amount of SiO₂ (CaCl₂ structure) added to the starting material (21). In addition to pv, another set of peaks not corresponding to any previously known phases appeared with particularly conspicuous peaks at 2.55 and 2.40 Å (marked H110 and H101, respectively). These two peaks are close to the diagnostic ppv peaks near 2.5 and 2.4 Å (ppv 202 and 110, respectively), but the high-quality XRD pattern clearly rejects the possibility of ppv. On the basis of three distinct new peaks at 4.433, 2.549, and 2.401 Å and several additional peaks (Tables 1 and 2), we concluded that this phase has a hexagonal unit cell of \( a = 5.096(2) \text{ Å} \) and \( c = 2.862(3) \text{ Å} \) and a molecular volume \( V = 32.19(4) \text{ Å}^3 \) with formula units per unit cell \( Z = 2 \) (Tables 1 and 2), referred to as the “H-phase.”

The powder XRD pattern was obtained with x-ray wavelength of 0.5164 Å.

**Fig. 1. In situ powder XRD patterns of Fs15 in Ne medium.** (A and B) Coexistence of pv phase and H-phase (A) during laser-heating at 2300 K and (B) after T quench to 300 K and 96 GPa. (C and D) Another coexisting pv phase and H-phase decompressed from 101 GPa (C) to 94 GPa at 300 K and then (D) to ambient conditions. Because of the multigrain spotty nature of the patterns, the relative peak intensities are affected by the presence or absence of strong spots. The peak at \( d \sim 2.7 \text{ Å} \), which is 78° in (B), belongs to a small amount of SiO₂ (CaCl₂ structure) added to the starting material (20).

**Table 1. Powder XRD data for the H-phase in Fs15 starting composition at 101 GPa.** The unit-cell parameters of the H-phase are \( a = 5.096(2) \text{ Å} \), \( c = 2.862(3) \text{ Å} \); \( \alpha = \beta = 90° \), \( \gamma = 120° \) at 101 GPa. The powder XRD pattern was obtained with x-ray wavelength of 0.5164 Å. \( h \), \( k \), \( l \) represent Miller indices of reflections; \( 2\theta \), Bragg angle; \( d \), \( d \)-spacing; obs, observed; calc, calculated; diff, difference.
The Bragg angle 2\(\theta\) was obtained for the H-phase with an x-ray wavelength of 0.3738 Å. The unit-cell parameters of the H-phase are calculated from the orientation matrix and observed as XRD spots.

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Fig. 2. Selection of MgSiO₃ pv and H-phase individual single crystals using the multigrain method. (A) A representative spotty multiple-crystal XRD pattern of the H-phase and coexisting pv in Fs15 at 101 GPa and after T quench (x-ray wavelength of 0.3738 Å), and the marked Miller indices show some of the most intense diffraction spots. (B) Six {101} diffraction reflections at different rotation angle \(\omega\) and azimuthal angle \(\eta\) belonging to one selected crystallite of the H-phase at 101 GPa.
together, indicating that they originate from the same single crystallite. We mapped the relatively large crystallites by using the x-ray beam, monitored the diffraction intensity, and observed that the intensities of the diffraction spots dropped ~50% after moving 0.5 μm in each direction, thus confirming the crystallite size to be typically ~0.5 μm or less (fig. S10).

**Fe Depletion in Perovskite**

All powder XRD peaks observed at 101 GPa could be assigned to the MgSiO₃ pv phase and H-phase. To confirm that the pv is indeed the MgSiO₃ endmember, we decompressed and monitored the disproportionation product with in situ XRD. The pv phase remains crystalline, with unit cell consistent with the MgSiO₃ pv all the way to ambient pressure (Fig. 1D), and all peaks in the powder XRD pattern can be assigned to the MgSiO₃ pv phase and H-phase. All powder XRD peaks observed at 101 GPa could be assigned to the MgSiO₃ pv phase and H-phase. We also performed scanning transmission electron microscopy (STEM) analysis on the recovered sample for phase identification and compositional analysis (20). First, we prepared a thin specimen with focused ion beam (FIB) and mapped it using energy-dispersive x-ray (EDX) to locate the MgSiO₃ region (Fig. 3). The STEM-EDX compositional mapping indicates four regions of distinct compositions marked as positions 1 to 4 on the TEM-bright field (BF) image (Fig. 3A). The selected area electron diffraction (SAED) in position 1 shows the quenched crystalline pv structure (Fig. 3G), and the chemical analysis verified $x = 0.02(2)$ based on multiple analyses (fig. S6); therefore, it can be nominally referred to as MgSiO₃. The STEM analysis shows that the composition and structure of the quenched pv match the XRD data of the same pv quenched to ambient conditions, confirming the Fe depletion in the pv as a result of the disproportionation. The overall quantity of SiO₂ in the recovered sample (position 2) is consistent with the excess SiO₂ in the starting composition and the in situ XRD observations (Fig. 3H and fig. S6). In contrast to the preservation of the crystalline MgSiO₃ pv, the H-phase was unquenchable and amorphized during decompression; its XRD peaks completely disappeared (Fig. 1D). The amorphous product was highly unstable under ambient conditions and further decomposed during the sample preparation process for STEM (positions 3 and 4). The Fe enrichment in H-phase was confirmed with EDX at ambient conditions (fig. S6), but its unit cell parameters and composition must be characterized in situ at high P. We also tested the reversibility of the disproportionation in the Ne medium by performing heating during decompression and observed that the H-phase clearly diminishes, and only the pv peaks remained at 67 GPa and 2000 K (fig. S7).

**Geophysical Implications**

The XRD pattern of the disproportionation products only contains the MgSiO₃ pv and H-phase, indicating that the H-phase has the opx stoichiometry and Fe enrichment ($x > 0.15$), as in the following reaction:

$$\text{Mg}_{0.85}\text{Fe}_{0.15}\text{SiO}_3 \rightarrow a\text{MgSiO}_3 + (1-a)\text{Mg}_{0.60}\text{Fe}_{0.40}\text{SiO}_3$$

According to the phase rule, at a given P-T condition compositions of both MgSiO₃ pv and the H-phase are fixed in the pseudo-binary system. The best approach to determine the H-phase composition, therefore, would be to synthesize a pure H-phase in situ. We tested another opx starting...
Cold compression of pv10 starting material in SiO2 glass medium and (Mg,Fe)SiO3 pv; no H-phase was observed.

assigned to three coexisting phases: SiO2, (Mg,Fe)Ofp, and (Mg,Fe)SiO3 pv. Cold compression of pv10 starting material in SiO2 glass medium and (Mg,Fe)SiO3 pv; no H-phase was observed.

The calculated peak positions of the H, pv, and SiO2 (CaCl2 structure) phases are indicated by small ticks, and the calculated peak positions of the ppv phase at the same P-T conditions are also shown for comparison.

Fig. 4. XRD powder patterns of run products from different starting compositions. (A) H-phase–dominant XRD pattern in the Fs40 composition at 85 GPa after T-quench to 300 K (Table 3). (B) Cold compression of pv10 starting material in SiO2 glass medium and (C) XRD pattern of coexisting pv phase and H-phase after T-quench from heating at 2350 to 300 K at 101 GPa. (D) XRD pattern of disproportionation products of (Mg0.80Fe0.20)(Al0.04Si0.96)O3 opx at 100 GPa after T-quench to 300 K from heating at 2300 K. The calculated peak positions of the H, pv, and SiO2 (CaCl2 structure) phases are indicated by small ticks, and the calculated peak positions of the ppv phase at the same P-T conditions are also shown for comparison.

material with the composition of (Mg0.80Fe0.20)SiO2 (Fs40), compressed it to 60 GPa, and laser heated it to 2100 K. Consistent with the previously known phase relations (25), all observed XRD peaks can be assigned to three coexisting phases: SiO2, (Mg,Fe)Ofp, and (Mg,Fe)SiO3 pv; no H-phase was observed. When the Fe40 was directly compressed to 85 GPa and heated to 2300 K, however, we observed a powder XRD pattern consisting of predominately H-phase, with only a minor amount of pv peaks (Fig. 4A), indicating that the H-phase composition is close to (Mg6.4Fe4.6)SiO3. Refinement of 15 powder XRD peaks of the H-phase yields a hexagonal unit cell of \( a = 5.136(6) \) Å, \( c = 2.883(3) \) Å, and \( V = 32.93(5) \) Å\(^3\) (Z = 2) for the H-phase at 85 GPa (Table 3). The volume reduction of the reaction based on the molar volumes of MgSiO3 pv and the H-phase on the right hand side and the pv on left side (21) was obtained as \( \Delta V = -0.27 \) Å\(^3\) and \( \Delta V/V = -0.2\% \) for the (Mg0.80Fe0.20)SiO3 H-phase. The small, negative \( \Delta V \) indicates that pressure slightly favors the disproportionation, thermodynamically.

Equation 1 implies that any pv in the composition range of \( x = 0.02 \) to 0.40 will disproportionate to MgSiO3 pv and H-phase at the deep lower mantle conditions. To test this conclusion, we used an (Mg0.60Fe0.40)SiO3 sample presynthesized in the pv form (pv10) with a multianvil press (Fig. 4B). The pv10 starting material is from the same batch that has been used extensively in previous studies and publications (6, 14, 24, 25) and is considered a realistic mineral component for the Earth’s interior. Again, we observed the disproportionation to an assemblage of Fe-depleted pv and H-phase (Fig. 4C). The direct transition from the stable phase of pv, rather than from a metastable opx starting material, demonstrates the relative stability. Because we can convert the Fe-bearing pv to coexisting Fe-depleted pv and H-phase, but have not been able to convert H-phase to Fe-bearing pv in this P-T range even with extensive heating, we conclude that the disproportionation assemblage is more stable than is Fe-bearing pv.

Most Earth models include excess fp in addition to the major mineral of pv. We conducted an experiment with the starting composition of MgO:FeO:SiO2 = 0.86:0.19:1.0 (pv:fp = 100:5) and applied directly for comparison with seismic data without adjustment for its Fe content. The sound-velocity data of (Mg,Fe)SiO3 pv may not be limited to the effects of the pv-ppv transition (4–6) in (Mg,Fe)SiO3 pv gives way to other Fe-bearing phases, such as oxides or H-phase, in this region. Likewise, interpretations for key enigmatic features in the deep lower mantle—namely, the sharp boundary on the top of the D” layer, the anticorrelations between \( V_{\text{S}} \) and \( V_{\text{P}} \), the seismic anisotropy, and the LLSPVs in the bottom third of the lower mantle (8–11, 27, 28)—may not be limited to the effects of the pv-ppv transition (7–3) but should consider the rich three-phase P-T-x relationship of pv, ppv, and the H-phase.

We still do not fully understand the detailed crystal chemistry of the H-phase. For instance,
the atomic positions have not been determined within its unit cell, and the small unit cell that we provided could well be the base for larger superlattices (20). Its composition and phase relation to pv and ppv needs to be determined as a function of P-T for interpreting various features in the deep lower mantle. Last, its elasticity and rheological properties are essential for interpretation of deep Earth seismology.

REFERENCES AND NOTES
20. Materials and methods are available as supplementary materials on Science Online.

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SUPPLEMENTARY MATERIALS
www.sciencemag.org/content/344/6186/sciencemag.org/suppl/DC1
Materials and Methods
Supplementary Text
Figs. S1 to S11
Tables S1 and S2
References (29–34)
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REPORTS
COSMOLOGY
Holographic description of a quantum black hole on a computer
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Black holes have been predicted to radiate particles and eventually evaporate, which has led to the information loss paradox and implies that the fundamental laws of quantum mechanics are strongly violated. A superstring theory, a consistent theory of quantum gravity, provides a possible solution to the paradox if evaporating black holes can actually be described in terms of standard quantum mechanical systems, as conjectured from the theory. Here, we test this conjecture by calculating the mass of a black hole in the corresponding quantum mechanical system numerically. Our results agree well with the prediction from gravity theory, including the leading quantum gravity correction. Our ability to simulate black holes offers the potential to further explore the yet mysterious nature of quantum gravity through well-established quantum mechanics.

In 1974, it was realized that a black hole should radiate particles as a perfect black body due to quantum effects in the surrounding space and that the black hole should eventually evaporate completely (1, 2). This phenomenon, which is known as the Hawking radiation, made more accurate the close analogy between the laws of black hole physics and those of thermodynamics (3), but it also caused a long scientific debate (4, 5) concerning the information-loss paradox (6, 7). Suppose one throws a book into a black hole. While the black hole evaporates, all we observe is the black-body radiation. Therefore, the information contained in the book is lost forever. This statement sharply conflicts with a basic consequence of the law of quantum mechanics that the information of the initial state should never disappear. Then the question is whether the law of quantum mechanics is violated or the above argument should somehow be modified if full quantum effects of gravity are taken into account.

Superstring theory, a consistent theory of quantum gravity, enabled us to understand the statistical-mechanical origin of the black hole entropy for a special class of stable black holes by counting the microscopic states, which look like the same black hole from a distant observer (8). The paradox still remains, however, because a complete description of an evaporating black hole is yet to be established. A key to resolve the paradox is provided by the gauge-gravity duality (9), which is conjectured in superstring theory. In a particular case, it claims that black holes in gravity theory can be described in terms of gauge theory in one dimension, which is a standard quantum mechanical system without gravity. The conjecture may be viewed as a concrete realization of the holographic principle (10, 11), which states that all of the information inside a black hole should be somehow encoded on its boundary, namely on the so-called event horizon of the black hole. In gauge theory in general, information is conserved during the time evolution as the standard laws of quantum mechanics simply apply. Therefore, it is widely believed that there is no information loss if the conjectured holographic description of a black hole is correct, including full quantum effects of gravity.

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